Assessing the uncertainty in permafrost lake margin detection from varying spatial resolution Synthetic Aperture Radar (SAR)

Anna Maria Trofaier¹
Gareth Rees¹, Barbara Widhalm², Annett Bartsch²

amat2@cam.ac.uk

 1 Scott Polar Research Institute, 2 Vienna University of Technology









VIENNA UNIVERSITY OF TECHNOLOGY DEPARTMENT OF GEODESY AND GEOINFORMATION

RESEARCH GROUPS
PHOTOGRAMMETRY & REMOTE SENSING

Why monitor permafrost lakes?

Why monitor permafrost?

- Permafrost is an indicator of climate change
- Thawing permafrost is a carbon source
- Transport in Arctic and sub-Arctic areas (roads, railways, pipelines) is affected by permafrost degradation
- Thawing of permafrost in alpine areas raises the risks of geohazards



Permafrost ground thermal regime changes due to:

- Changes in air temperature and/or
- Surface disturbances
- Precipitation
- Clearing of vegetation
- Removal of insulating organic layer
- Forest fires
- River channel migration
- Shoreline erosion

Response of Permafrost to climate change depends on variations in local seasonal factors:

- Snow cover
- Vegetation
- Surficial material
- Moisture content
- Drainage

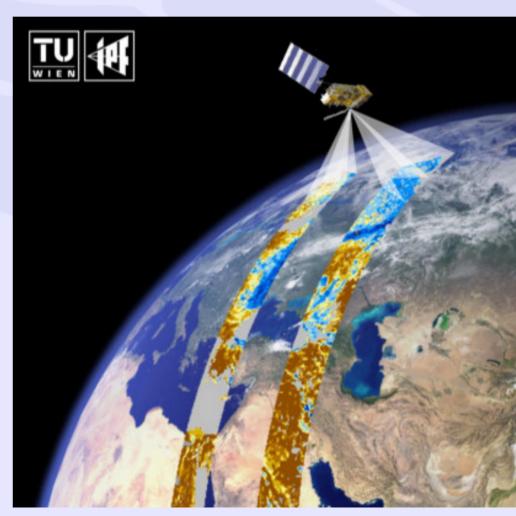


Permafrost & Remote Sensing

Cannot directly see below the soil surface, but need to monitor indicators:

- Land-surface hydrology
- Terrain changes
- Vegetation





Land-surface hydrology

- Surface soil moisture (SSM)
- -soil moisture influences heat transfer
- -Indicative for drainage conditions
- Lakes (water bodies WB)
- –Indicative for thermokarst processes, permafrost degradation

- State transition (freeze/thaw timing – FT; surface state flag - SSF)
- -Complements land surface temperature
- -Can serve as masking input for SSM



Related Work: Smith et al. 2005 in Science

BREVIA

Disappearing Arctic Lakes

L. C. Smith, 1* Y. Sheng, 2 G. M. MacDonald, 1 L. D. Hinzman 3

Arctic warming has accelerated since the 1980s, driving an array of complex physical and ecological changes in the region (1). Particularly puzzling has been evidence for perturbations to the terrestrial water cycle (2), which plays an integral role in nearly every aspect of the Arctic system. We compared

satellite imagery acquired across ~515,000 km2 of Siberia in the early 1970s with recent (1997 to 2004) satellite data to inventory and track ongoing changes in more than 10,000 large lakes after three decades of rising soil and air temperatures in the region (1, 3, 4). Our analysis reveals a widespread decline in lake abundance and area, despite slight precipitation increases (4). The spatial pattern of lake disappearance strongly suggests that thawing of permafrost is driving the observed losses.

Between 1973 and 1997-98, the total number of large lakes (those >40 ha) decreased from 10,882 to 9712, a decline of 1170 or ~11% (SOM text). Most did not disappear altogether, but instead shrank to sizes below 40 ha. Total regional lake surface area decreased by 93,000 ha, a ~6% decline. One hundred and twenty-five lakes vanished completely and are now revegetated, as indicated by sharp increases in near-infrared reflectance (Fig. 1, B and C). Sub-

ronments, driven primarily by slumping and collapsed terrain features (thermokarst) that subsequently fill with water (SOM text). Such observations are in apparent conflict with the phenomenon seen here and also near Council, Alaska, where thermokarst ponds in discontinuous permafrost are also shrinking (5).

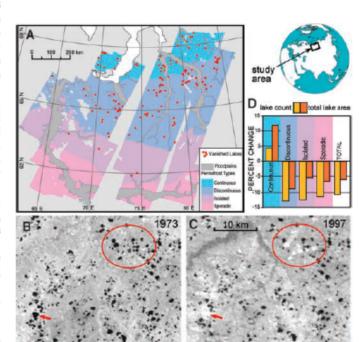


Fig. 1. (A) Locations of Siberian lake inventories, permafrost distribution, and vanished lakes. Total lake abundance and inundation area have declined since 1973 (B), including (C) permanent drainage and revegetation of former lakebeds (the arrow and oval show representative areas). (D) Net

rather than a direct climatic mechanism such as increased evaporation). It also raises the possibility of a diffuse lake drainage "front" where warming permafrost first experiences widespread degradation. The fact that ~85% of the vanished lakes reported here occur within 200 km of the continuous permafrost boundary lends some support to this concept (Fig. 1A).

Clearly, other factors besides permafrost influence substrate permeability and lake drainage. In west Siberia, shallow water tables and extensive, low-permeability peatlands (6) en-

> sure continued survival of many lakes, even where permafrost is absent. Overlay of our lake maps with a detailed peatland inventory (7) shows that, although lakes in continuous permafrost are found on all substrates, they exist only as perched systems on peatlands further south. In such regions, factors besides permafrost degradation will be important to lake persistence. However, aside from lowpermeability environments and/or beneficent water balance adjustments (i.e., further increases in net precipitation), the ultimate effect of continued climate warming on high-latitude, permafrost-controlled lakes and wetlands may well be their widespread disappearance.

References and Notes

- 1. M. C. Serreze et al., Clim. Change 46,
- B. J. Peterson et al., Science 298, 2171 (2002).
- 3. A. V. Pavlov, N. G. Moskalenko, Permafrost Periglac. Process. 13, 43 (2002).
- 4. K. E. Frey et al., Polar Res. 22, 287
- 5. K. Yoshikawa, L. D. Hinzman, Permafrost Periglac. Process. 14, 151 (2003).

Related Work: Smith et al. 2005 in Science

Karlsson et al. 2012 in Journal of Hydrology

BREVIA

Disappearing Arctic Lakes

L. C. Smith, 1* Y. Sheng, 2 G. M. MacDonald, 1 L. D. Hinzman 3

rather than a direct climatic mechanism such as increased evaporation). It also raises the possibility of a diffuse lake drainage "front" where warming permafrost first experiences widespread degradation. The fact that ~85% of the

within 200

HYDROLOGY

Journal of Hydrology 464-465 (2012) 459-466



Journal of Hydrology

journal homepage: www.elsevier.com/locate/jhydrol



Thermokarst lake, hydrological flow and water balance indicators of permafrost change in Western Siberia

Johanna Mård Karlsson*, Steve W. Lyon¹, Georgia Destouni²

Department of Physical Geography & Quaternary Geology, Stockholm University, SE-106 91 Stockholm, Sweden Bert Bolin Centre for Climate Research, Stockholm University, SE-106 91 Stockholm, Sweden

ARTICLE INFO

Article history: Received 15 February 2012 Received in revised form 5 July 2012 Accepted 21 July 2012 Available online 31 July 2012 This manuscript was handled by Konstantine P. Georgakakos, Editor-in-Chief, with the assistance of Matthew Rodell, Associate Editor

Keywords: Per mafrost Thermokarst lakes Hydrological cycles and budgets Catchments Climate change Water storage change

SUMMARY

Permafrost, mainly of discontinuous type, that underlies the tundra and taiga landscapes of the Nadym and Pur river basins in northwestern Siberia has been warming during the recent decades. A mosaic of thermokarst lakes and wetlands dominates this area. In this study we tested the hypothesis chain that permafrost thawing changes thermokarst lake area and number, and is then also reflected in and detectable through other associated hydrological changes. Based on indications from previous studies, the other hydrological changes in a basin were expected to be decreasing intra-annual runoff variability (quantified by decreasing maximum and increasing minimum runoff) and systematically decreasing water storage. To test this hypothesis chain, we mapped thermokarst lake changes using remote sensing analysis and analyzed both climate (temperature and precipitation) and water flow and balance changes using available monthly data records. This was done for the whole Nadym and Pur river basins and a smaller subbasin of the former (denoted 7129) with comparable data availability as the whole river basins. The results for the 7129 sub-basin show all the indicators (thermokarst lake and other hydrological) changing consistently, as could be expected in response to permafrost thawing that alters the connections between surface and subsurface waters, and leads to overall decreases in water (including ground ice) storage within a basin. Over the Nadym and Pur basins, the relative area influenced by similar permafrost thawing and associated lake and hydrological effects appears (yet) too small to be clearly and systematically reflected in the basin-average indicators for these large basins.

Change 46, e 298, 2171

dary lends

ake drain-

tables and

ds (6) enof many

nafrost is

ake maps

inventory n lakes in are found

ist only as

tlands furns, factors

gradation

ce persist-

from lownts and/or ce adjust-

ases in net

ate effect

arming on

controlled

v well be

earance.

A). nafrost in-

enko, Perma-43 (2002). es. 22, 287

n, Permafrost 2003).

- Research: Monitor extent and dynamics of spring floods
- Motivation: Inter-annual changes in water body extent remain inconspicuous.

1498

JOURNAL OF HYDROMETEOROLOGY

VOLUME 12

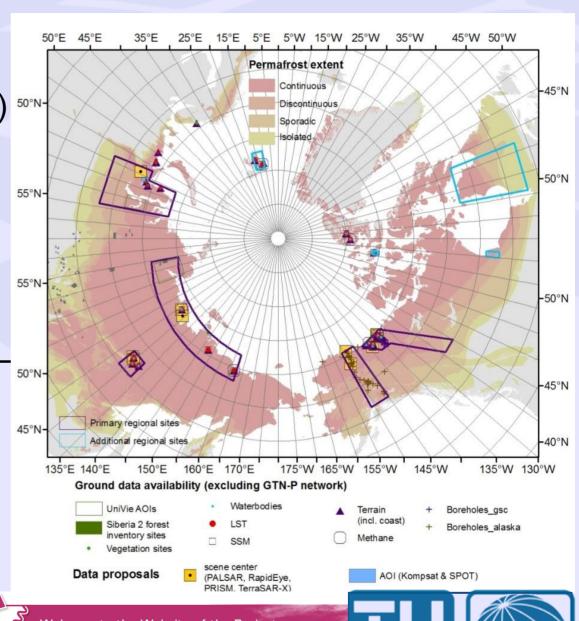
"Spring flood can start and end at different dates and its duration may vary significantly from one year to another, and thus some of the interannual variability of snowmelt is hidden when monthly or annual averages are used." Zakharova et al. (2011)

Snow Cover and Spring Flood Flow in the Northern Part of Western Siberia (the Poluy, Nadym, Pur, and Taz Rivers)

E. A. ZAKHAROVA,*,*,* A. V. KOURAEV,*,* S. BIANCAMARIA,*,*,* M. V. KOLMAKOVA,*,[®] N. M. MOGNARD,*,[&] V. A. ZEMTSOV,[®] S. N. KIRPOTIN,[®] AND B. DECHARME**

Regional lake monitoring

- ■Medium scale monitoring is possible by use of SAR (radar) 50°Ndata.
- Fairly good coverage: ENVISAT ASAR Wide Swath:
- -120 m resolution
- -Open water surfaces detectable
- Annual maps of open water/ water fraction (summer stage -July-August) since 2007 + number of acquisitions





Welcome to the Website of the Project

DUE PERMAFROST

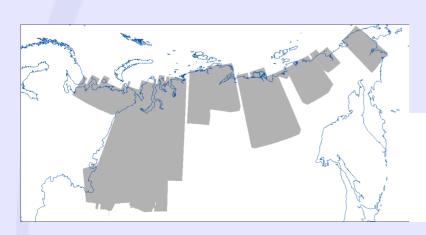


Regional water bodies

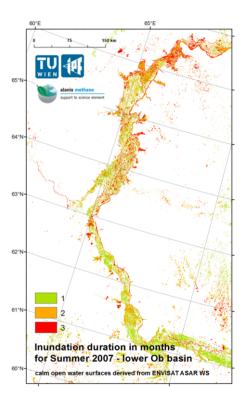
Seasonal dynamics for Northern Eurasia will become

available from ESA STSE ALANIS

- www.alanis-methane.info
- Specification
 - 10 day steps with update flag
 - 2007 and 2008
 - Coverage: parts of northern Eurasia









'Local' wetlands product



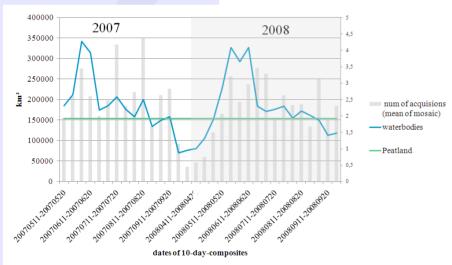
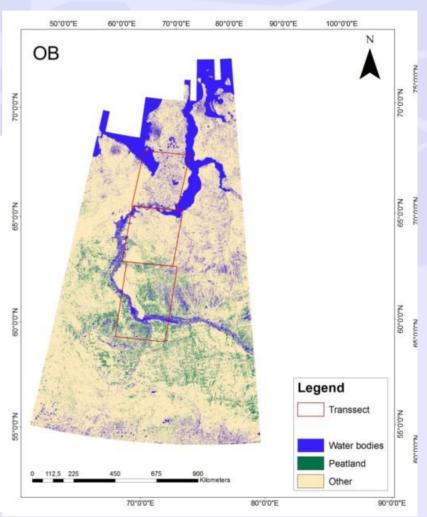
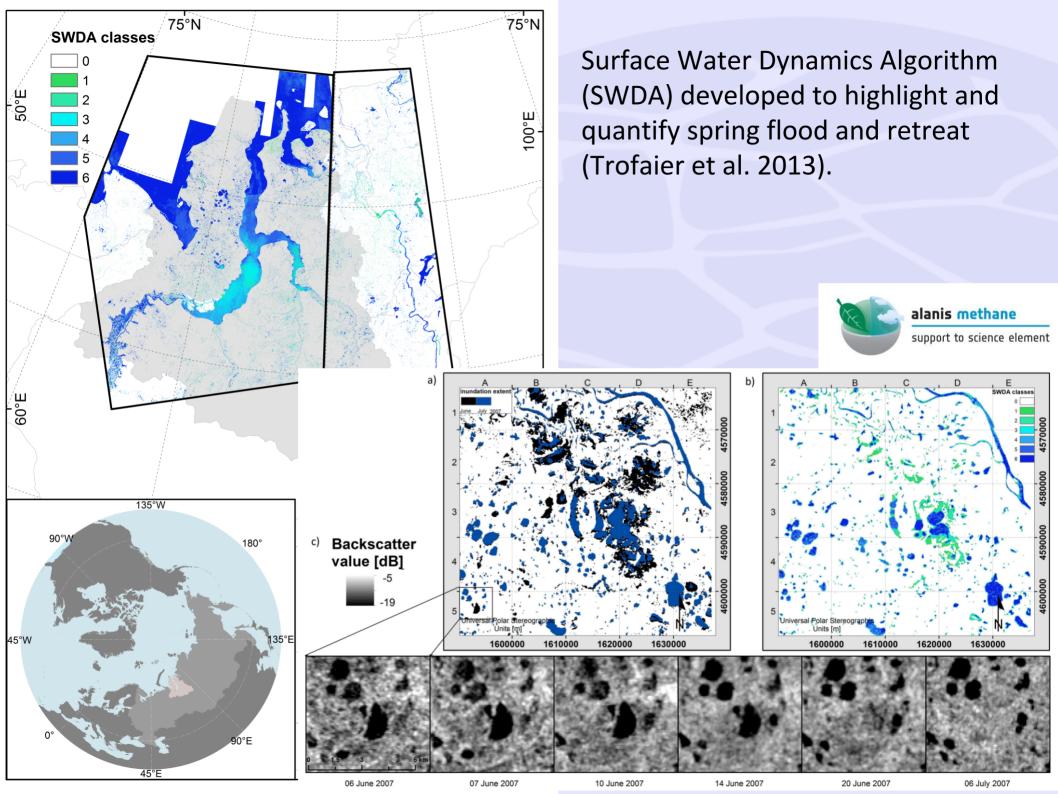


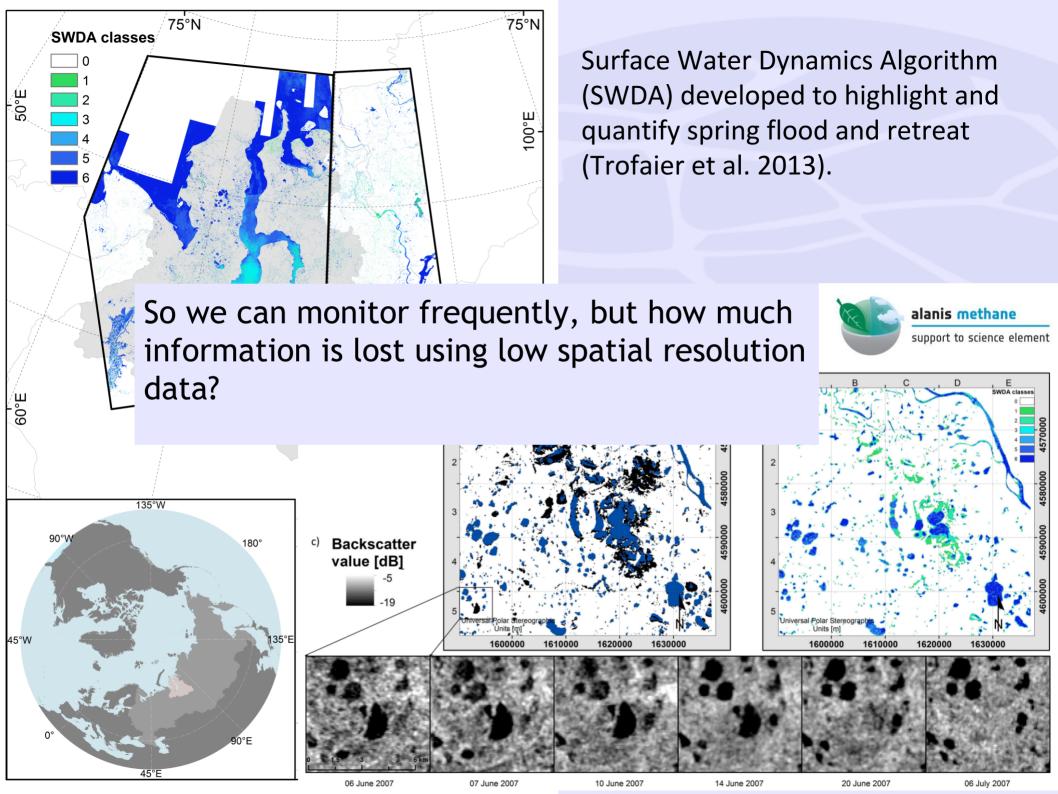
Figure 7: Temporal inundation dynamics of water bodies (blue) in relation with peatland area (green) in a **Ob river** basin and delta derived from ASAR WS data of 2007 and 2008.











Yamal Peninsula

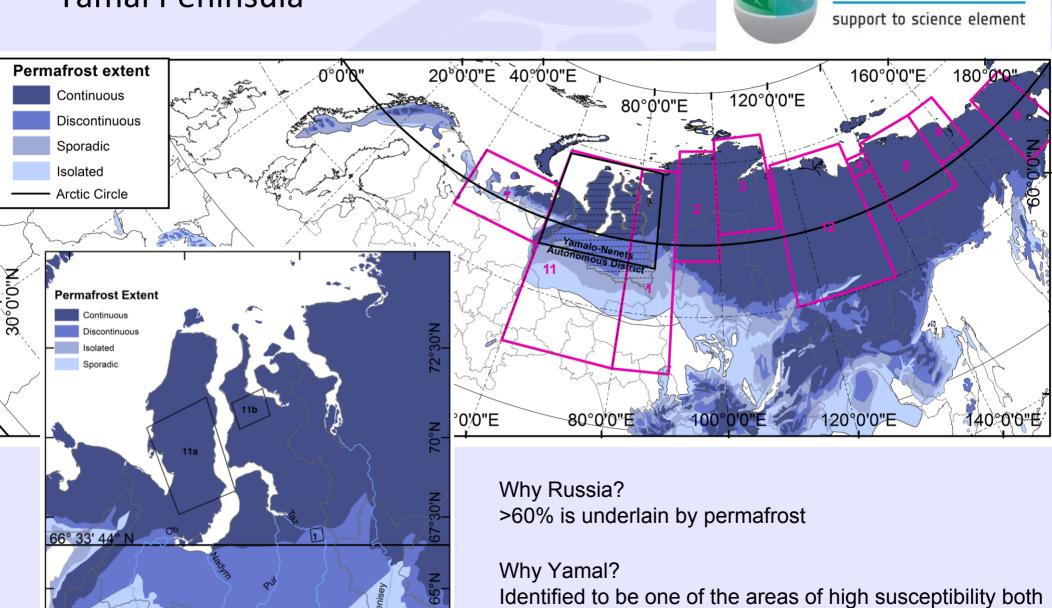
80°E

70°E

90°É



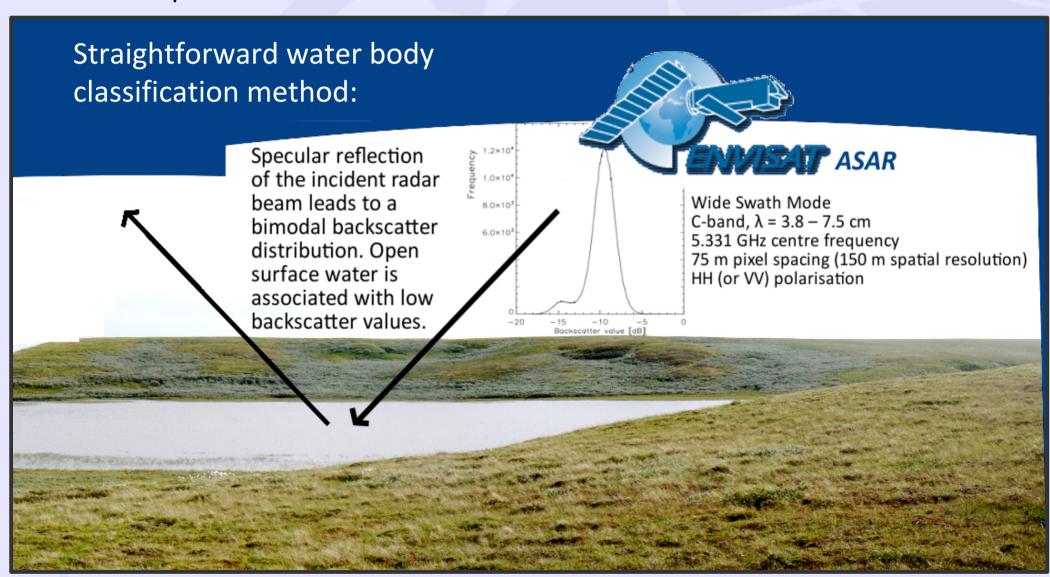
to climate-induced changes and human activities



Active Microwave Data:

Ideally suited for monitoring surface hydrology.

- -Active: Independent of solar illumination
- -Cloud penetration



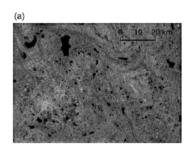
Thresholding:

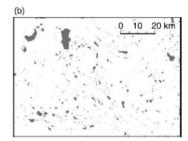
Classification of water bodies through thresholding. Bartsch et al. 2008:

89 © IWA Publishing 2008 Hydrology Research 39.2 2008

Detection of permanent open water surfaces in central Siberia with ENVISAT ASAR wide swath data with special emphasis on the estimation of methane fluxes from tundra wetlands

Annett Bartsch, Carsten Pathe, Klaus Scipal and Wolfgang Wagner





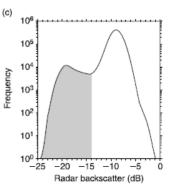


Figure 3 | Threshold classification example of a tundra site: (a) grayscale normalized image, (b) classified image with lakes in gray and (c) histogram of normalized backscatter in dB

Lakes from ENVISAT ASAR WS - Known problems (i)

- C-Band sensitivity to weather in case of this specific application
- Number of acquisitions – all available data independent from weather condition!
- Up to 60% of data affected
- Investigated within ESA STSE ALANIS Methane

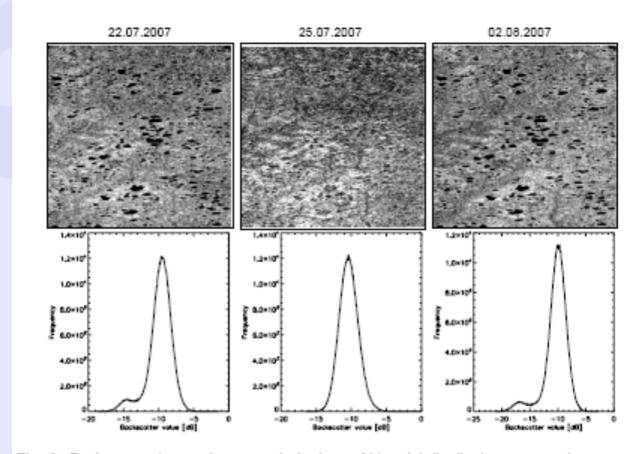


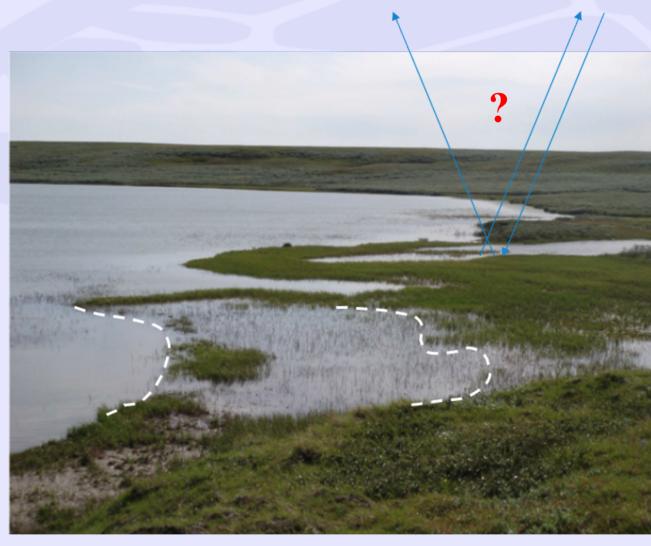
Fig. 3. Backscatter time series example for loss of bimodal distribution over tundra: top normalized backscatter images, bottom: histograms of backscatter distribution [dB].

Bartsch et al. 2012



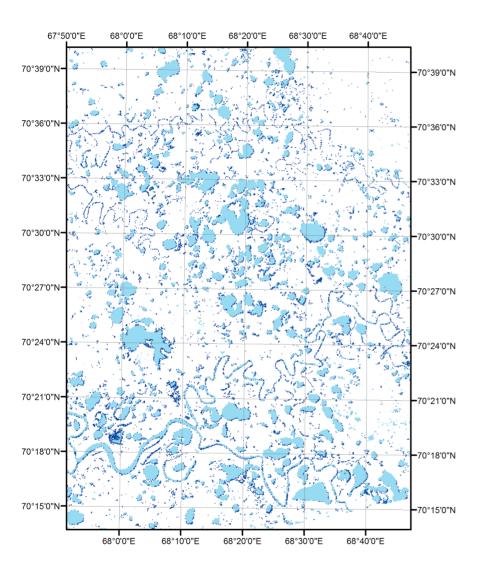
Lakes from ENVISAT ASAR WS - Known problems (ii)

Radar – sensitivity (in particular C-Band) to vegetation results in double bounce rather than specular reflection.



Trofaier et al. 2013

Assessment of wavelength impact and lake size



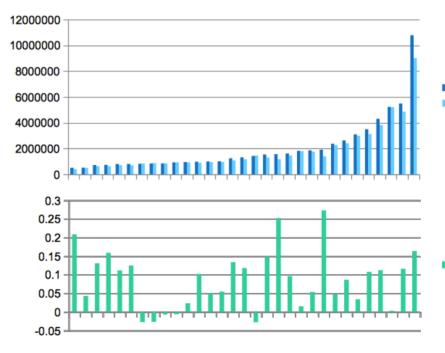
water bodies

ASAR WS (15.07.2007)

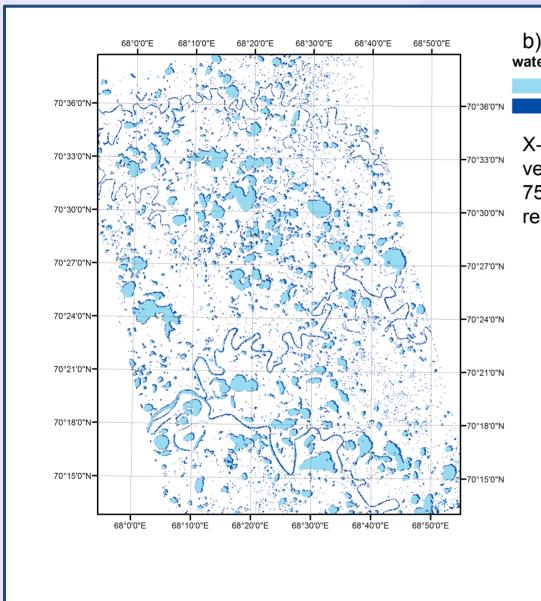
ALOS PALSAR (14.07.2007)

L-band (HV, 16 m) versus C-band (VV, 75 m nominal resolution) PALSAR vs. ASAR WS

Comparison of ALOS PALSAR and ASAR WS open water extent by lake size



Assessment of wavelength impact and lake size



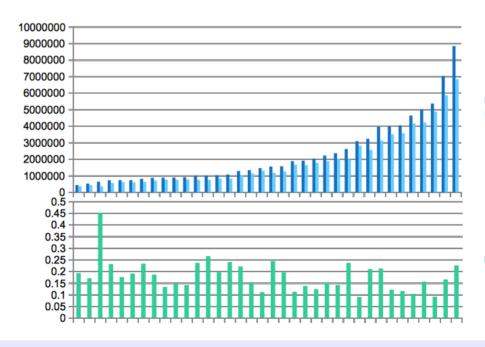


ASAR WS (13.08.2008)
TSX (13.08.2008)

X-band (HH, 5m) versus C-band (VV, 75 m nominal resolution)

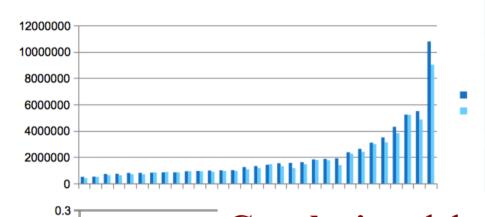


Comparison of TerraSAR-X and ASAR WS open water extent by lake size



Assessment of wavelength impact and lake size

Comparison of ALOS PALSAR and ASAR WS open water extent by lake size

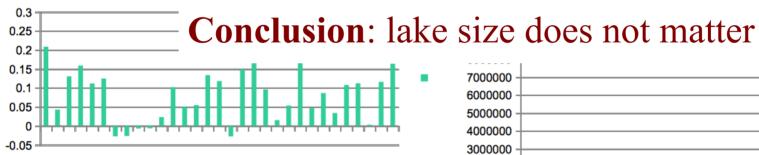


b)

X-band (HH, 5m) versus C-band (VV, 75 m nominal resolution)

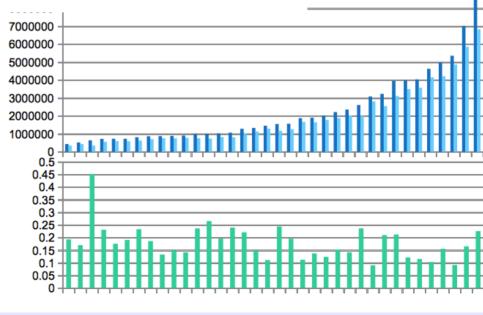
TerraSAR-X vs. ASAR WS

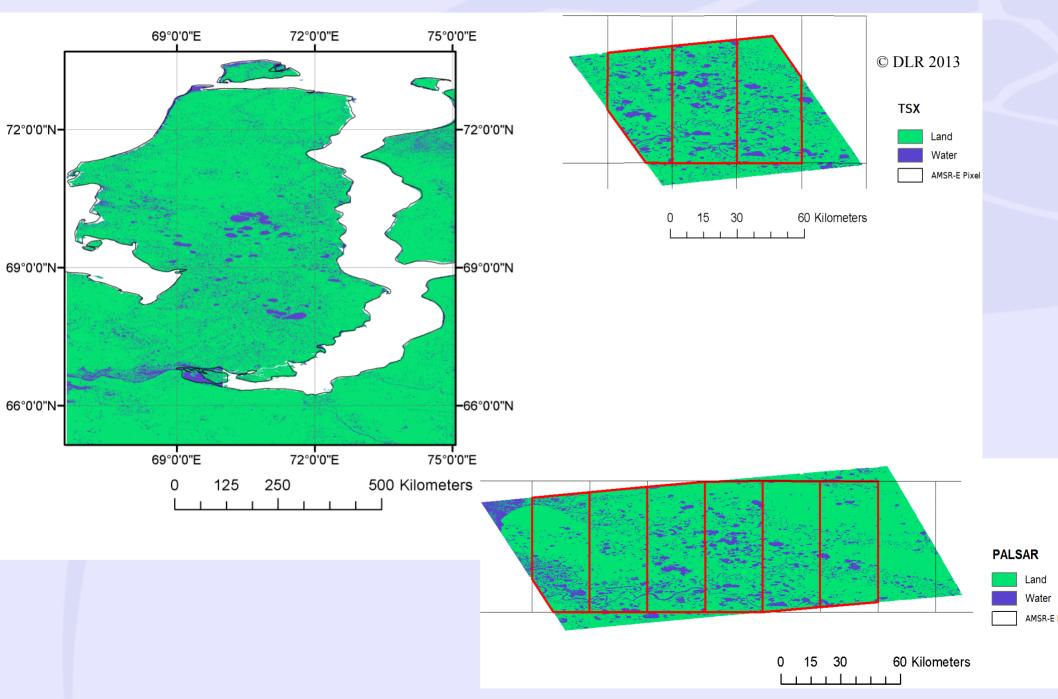
Comparison of TerraSAR-X and ASAR WS open water extent by lake size

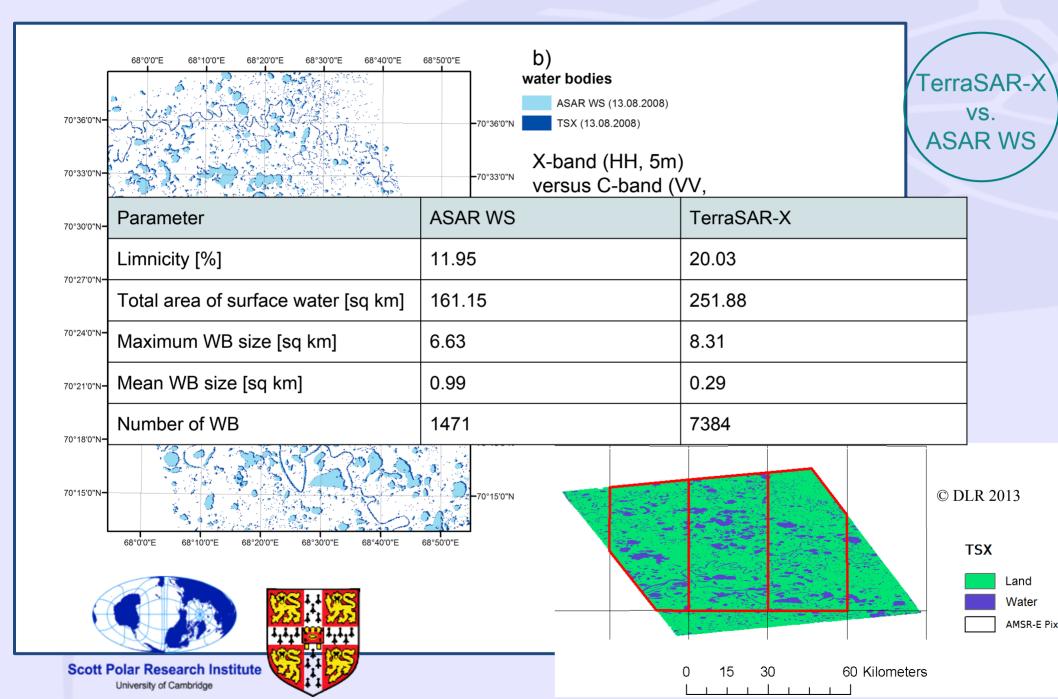


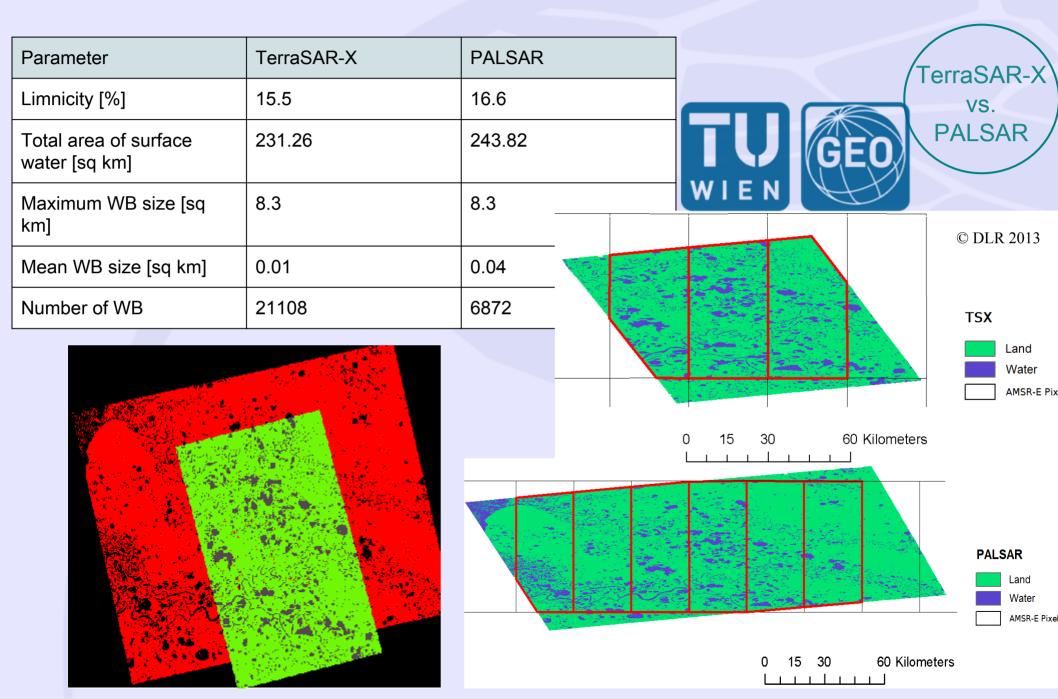
palsar - Asar [%]
vs. I -band (HV 16 m) versus

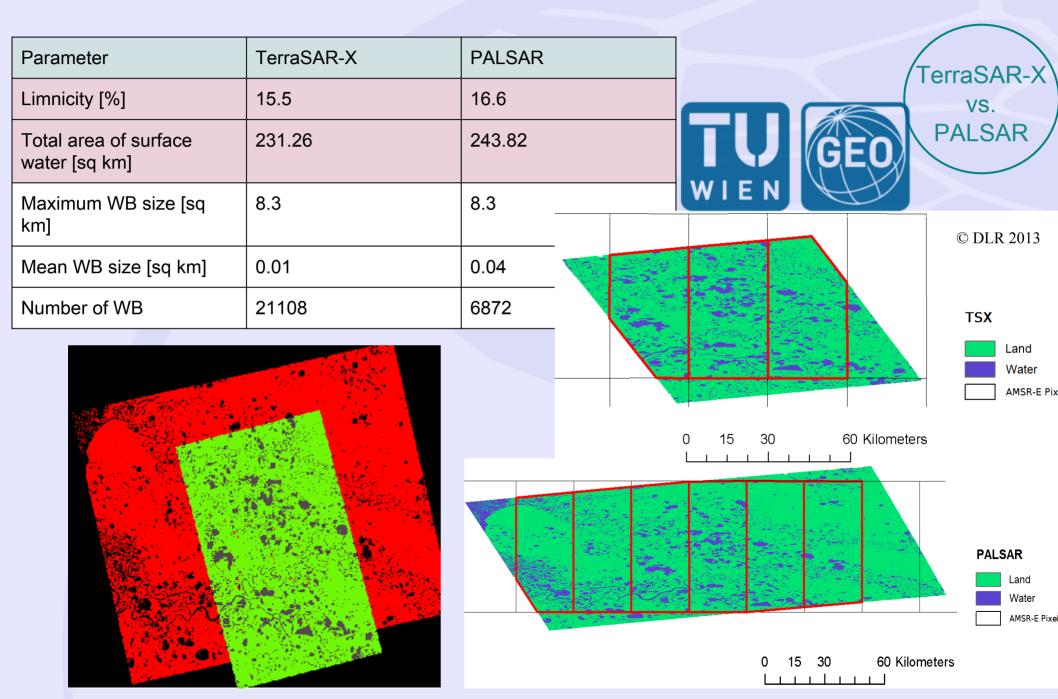
VS. L-band (HV, 16 m) versus
ASAR WS C-band (VV, 75 m
nominal resolution)



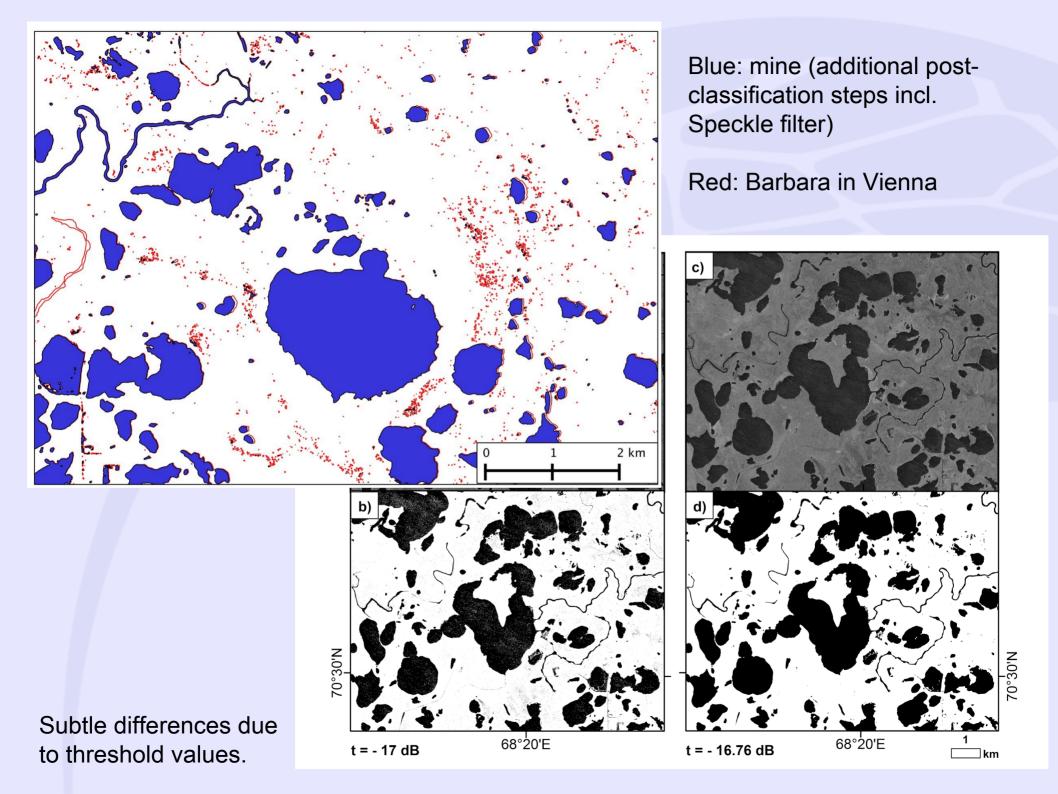








| Parameter | TerraSAR-X TU WIEN GEO | TerraSAR-X Scott Polar Research Institute University of Cambridge |
|-------------------------------------|---------------------------|--|
| Limnicity [%] | 15.5 | 20.03 TU are looking |
| Total area of surface water [sq km] | 231.26 | 251.88 at smaller area to compare to PALSAR |
| Maximum WB size [sq km] | 8.3 | 8.31 |
| Mean WB size [sq km] | 0.01 | 0.29 |
| Number of WB | 21108 | 7384 |



| Parameter | TerraSAR-X | PALSAR |
|-------------------------------------|------------|--------|
| Limnicity [%] | 15.5 | 16.6 |
| Total area of surface water [sq km] | 231.26 | 243.82 |
| Maximum WB size [sq km] | 8.3 | 8.3 |
| Mean WB size [sq km] | 0.01 | 0.04 |
| Number of WB | 21108 | 6872 |

How come TerraSAR-X captures less surface water than PALSAR?

TerraSAR-X: X-band (HH, 5m spatial resolution)

ALOS PALSAR: L-band (HV, 16 m spatial resolution)

| Parameter | TerraSAR-X | PALSAR |
|-------------------------------------|------------|--------|
| Limnicity [%] | 15.5 | 16.6 |
| Total area of surface water [sq km] | 231.26 | 243.82 |
| Maximum WB size [sq km] | 8.3 | 8.3 |
| Mean WB size [sq km] | 0.01 | 0.04 |
| Number of WB | 21108 | 6872 |

How come TerraSAR-X captures less surface water than PALSAR?

TerraSAR-X: X-band (HH, 5m spatial resolution)
ALOS PALSAR: L-band (HV, 16 m spatial resolution)



Conclusions

Lake monitoring is possible at low spatial resolution AND is needed to monitor seasonal lake and wetland dynamics - example Yamal peninsula:

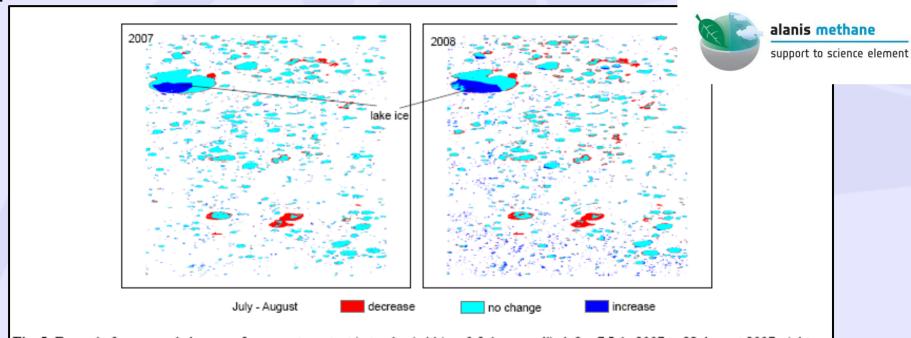


Fig. 5. Example for seasonal changes of open water extent in tundra (within a 0.5 degree cell): left – 7 July 2007 to 28 August 2007, right – 6 July 2008 to 21 August 2008.

Bartsch et al., 2012 Trofaier et al. 2013

HOWEVER, need to be careful when classifying lakes due to emergent vegetation over the growing season.